

# Water Vapour Feedback

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In the climate debate, the argument of feedback through water vapor is used to amplify the climate effect of greenhouse gases – the sensitivity to a doubling of their concentration in the atmosphere – which, according to the radiative transfer equation and general consensus, is a maximum of  $0.8^{\circ}$ , by an alleged factor of 2-6. However, this is usually not quantified more precisely, only formulas with the „final feedback“ are usually given.

Recently, David Coe, Walter Fabinski and Gerhard Wiegler described and analyzed precisely this feedback in the publication „[The Impact of CO<sub>2</sub>, H<sub>2</sub>O and Other „Greenhouse Gases‘ on Equilibrium Earth Temperatures](#)„. Based on her publication, this effect is derived below using partly the same and partly slightly different approaches. The results are almost identical.

All other effects that occur during the formation of water vapor, such as cloud formation, are ignored here.

## The basic mechanism of water vapor feedback

The starting point is an increase in atmospheric temperature by  $\Delta T_0$ , regardless of the cause. Typically, the greenhouse effect is assumed to be the primary cause. The argument is now that the warmed atmosphere can absorb more water vapor, i.e. the saturation vapor pressure (SVP) increases and it is assumed that consequently the water vapor concentration  $\Delta H_{20}$  also increases, as a linear function of the temperature change. (The temperature change is so small that linearization is legitimate in any case):

$$\Delta H_{20} = j \cdot \Delta T_0$$

where  $j$  is the proportionality constant for the water vapor

concentration.

An increased water vapor concentration in turn causes a temperature increase due to the greenhouse effect of water vapor, which is linearly dependent on the water vapor concentration:

$$\Delta T_1 = k \cdot \Delta H_{20}$$

In summary, the triggering temperature increase  $\Delta T_0$  causes a subsequent increase in temperature  $\Delta T_1$ :

$$\Delta T_1 = j \cdot k \cdot \Delta T_0$$

Since the prerequisite of the method is that the cause of the triggering temperature increase is insignificant, the increase by  $\Delta T_1$  naturally also causes a feedback cycle again:

$$\Delta T_2 = j \cdot k \cdot \Delta T_1 = (j \cdot k)^2 \cdot \Delta T_0$$

This is repeated recursively. The final temperature change is therefore a geometric series:

$$\Delta T = \Delta T_0 \sum_{n=0}^{\infty} (j \cdot k)^n = \Delta T_0 \cdot \frac{1}{1 - j \cdot k}$$

If  $j \cdot k \geq 1$ , the series would diverge and the temperature would grow beyond all limits. It is therefore important to be clear about the magnitude of these two feedback factors.

For the determination of the first term,  $j$  we can apply a simplified approach by accepting the statement commonly used in the mainstream literature, that for each degree C of temperature increase the relative air moisture may rise up to 7%. [In the German version of this post](#) I did the explicit calculations and came to the result that the realistic maximum air moisture rise is 6% per degree temperature rise, which has hardly any effect on the final result.

## **Dependence of the greenhouse effect on the change in relative humidity**

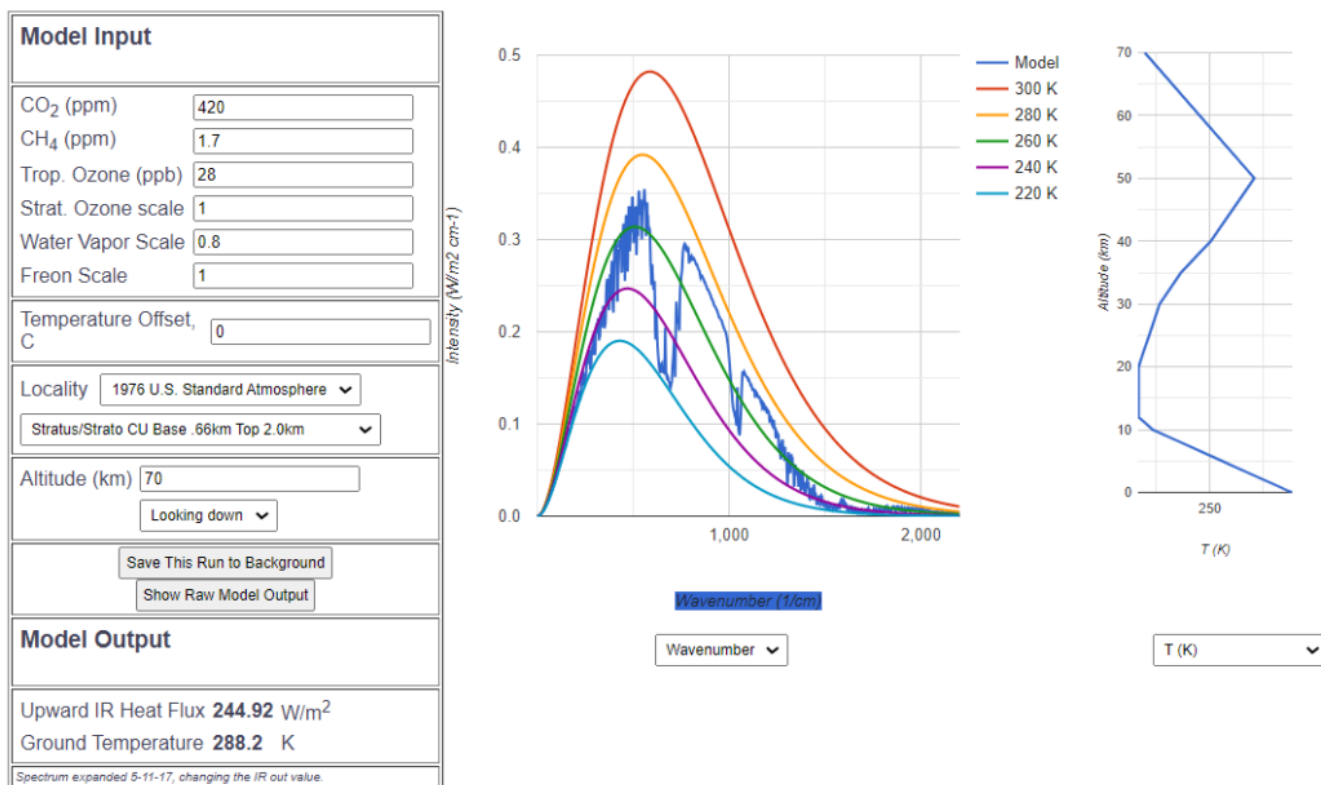
Infrared radiation transport in the atmosphere is dependent on relative humidity. This is taken into account in the well-known and proven MODTRAN simulation program. With increasing

humidity, the outgoing infrared radiation decreases due to the greenhouse effect of water vapor.

The decrease in radiation is linear between 60% and 100% humidity. Therefore, the increase in relative humidity from 80% to 86% is considered to determine the decrease in radiant power and the temperature increase required for compensation.

To do this, we set the parameters of the MODTRAN simulation to

- the current CO<sub>2</sub> concentration of 420 ppm,
- a relative humidity of 80%,
- and a cloud constellation that comes close to the average IR radiant power of  $240 \frac{W}{m^2}$ .



The temperature offset is now increased until the reduced iR radiation of  $0.7 \frac{W}{m^2}$  is compensated for by increasing the temperature. This is the case when the ground temperature is increased by 0.215 °C.

Model Input		Model Input	
CO <sub>2</sub> (ppm)	420	CO <sub>2</sub> (ppm)	420
CH <sub>4</sub> (ppm)	1.7	CH <sub>4</sub> (ppm)	1.7
Trop. Ozone (ppb)	28	Trop. Ozone (ppb)	28
Strat. Ozone scale	1	Strat. Ozone scale	1
Water Vapor Scale	0.87	Water Vapor Scale	0.87
Freon Scale	1	Freon Scale	1
Temperature Offset, C	0	Temperature Offset, C	0.215
Locality	1976 U.S. Standard Atmosphere	Locality	1976 U.S. Standard Atmosphere
Stratus/Strato CU Base .66km Top 2.0km		Stratus/Strato CU Base .66km Top 2.0km	
Altitude (km)	70	Altitude (km)	70
Looking down		Looking down	
Delete Background Model Run		Delete Background Model Run	
Show Raw Model Output		Show Raw Model Output	
Model Output		Model Output	
Upward IR Heat Flux	244.104 W/m <sup>2</sup>	Upward IR Heat Flux	244.92 W/m <sup>2</sup>
IR Heat Loss (Background)	244.92 W/m <sup>2</sup>	IR Heat Loss (Background)	244.92 W/m <sup>2</sup>
... Difference, New - BG	-0.82 W/m <sup>2</sup>	... Difference, New - BG	0 W/m <sup>2</sup>
Ground Temperature	288.2 K	Ground Temperature	288.42 K
Spectrum expanded 5-11-17, changing the IR out value.		Spectrum expanded 5-11-17, changing the IR out value.	

A 7% higher relative humidity therefore causes a greenhouse effect, which is offset by a temperature increase of 0.215°C. Extrapolated to a (theoretical) change of 100% humidity, this results in  $k=3.07^\circ\text{C}/100\%$ .

## The final feedback factor and the total greenhouse effect

This means that a 1 degree higher temperature in a feedback cycle causes an additional temperature increase of  $k \cdot j = 0.215^\circ\text{C}$ .

The geometric series leads to an amplification factor  $f$  of

the pure CO<sub>2</sub> greenhouse effect by

$$f = \frac{1}{1-0.215} = 1.27$$

This means that the sensitivity amplified by the water vapor feedback when doubling the CO<sub>2</sub> concentration  $\Delta T$  is no longer  $\Delta T_0 = 0.8^\circ\text{C}$ , but

$$\Delta T = 1.27 \cdot 0.8^\circ\text{C} = 1.02^\circ\text{C} \approx 1^\circ\text{C}$$

This result does not take into account the increase in temperature caused by the higher water vapor concentration